

# GEOLOGIC MAP OF THE DOLAN SPRINGS 7 1/2 QUADRANGLE, MOHAVE COUNTY, ARIZONA

by  
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Arizona Geological Survey Digital Geologic Map 81  
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Geologic map of the Dolan Springs 7 1/2 Quadrangle,  
Mohave County, Arizona: Arizona Geological Survey  
Digital Geologic Map DGM-81, version 1.0, scale 1:24,000.

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## Introduction

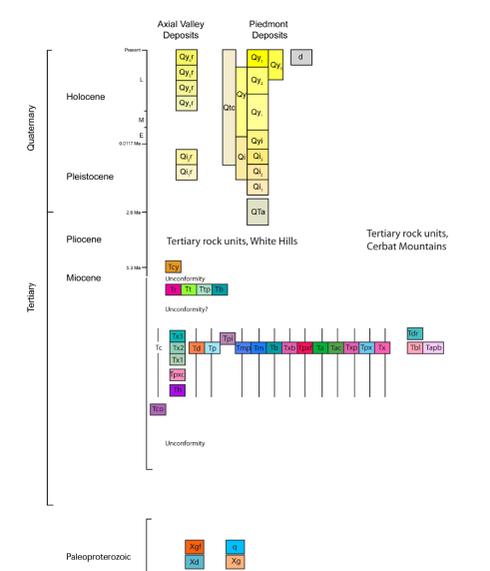
The Dolan Springs 7 1/2 Quadrangle encompasses part of Detrital Valley, the southernmost  
White Hills, and the northernmost Cerbat Mountains in northwestern Arizona. This map depicts  
the Precambrian and Cenozoic bedrock in the mountains and late Cenozoic surficial deposits in  
the valley. Production of this new geologic map continues the Arizona Geological Survey  
mapping program in northwestern Arizona that includes the Grasshopper Junction 7 1/2  
Quadrangle immediately to the south (Ferguson et al., 2009) and several 7 1/2 quadrangles  
farther southwest in the Colorado River Valley. This mapping was done under the joint State  
Federal STATEMAP program, as specified in the National Geologic Mapping Act of 1992, and  
was jointly funded by the Arizona Geological Survey and the U.S. Geological Survey under  
STATEMAP Program Contract award number 08HQAG0093. Mapping was compiled digitally  
using ESRI ArcGIS software.

The map area is located in Detrital Valley in the Basin and Range Province of northwestern  
Arizona. Most of the Dolan Springs Quadrangle is underlain by late Cenozoic valley-filling  
clastic sedimentary deposits that are as much as 300 feet thick (Richard et al., 2007; Mason et  
al., 2007). These consist mostly of sandstone and conglomerate in alluvial fans that flank the  
valley and a fine-grained axial facies that locally includes halite (Mason et al., 2007). The  
surface of the valley is covered by relatively thin deposits of various ages ranging from late  
Pliocene to modern. In dramatic contrast to the Colorado River Valley to the west, Detrital Valley  
has undergone very little incision during the Quaternary, and thus the valley topography is  
relatively smooth. In general, upper piedmont areas are dominated by moderately coarse  
Pliocene alluvial fan deposits, and young, relatively fine-grained deposits are prevalent on  
the lower piedmonts and along Detrital Wash. Extensive, complex distributary drainage systems  
in middle piedmont areas are outlined by the extent of late Holocene deposits. In these areas,  
this geologic map indicates areas of potential flood hazards due to complex flow paths and  
widespread sheetflooding.

Bedrock units of three types are present: Proterozoic metamorphic and plutonic rocks, and two  
suites of Miocene volcanics. Most of the volcanics are intermediate to mafic lavas with  
associated breccia and conglomeratic rocks. These rocks overlie crystalline basement along a  
fairly steeply west-dipping unconformity at the western foot of the Cerbat Range, and again in  
one area just north of Dolan Springs where the strata are gently east-dipping. A major down-  
throw normal fault partially separates the Cerbats from the White Hills. Intermediate to  
mafic lavas are complexly interbedded and display no obvious stratigraphic sequence or regional  
markings. Stratigraphic relationships change abruptly along strike and between adjacent fault  
blocks.

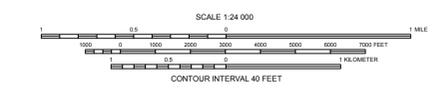
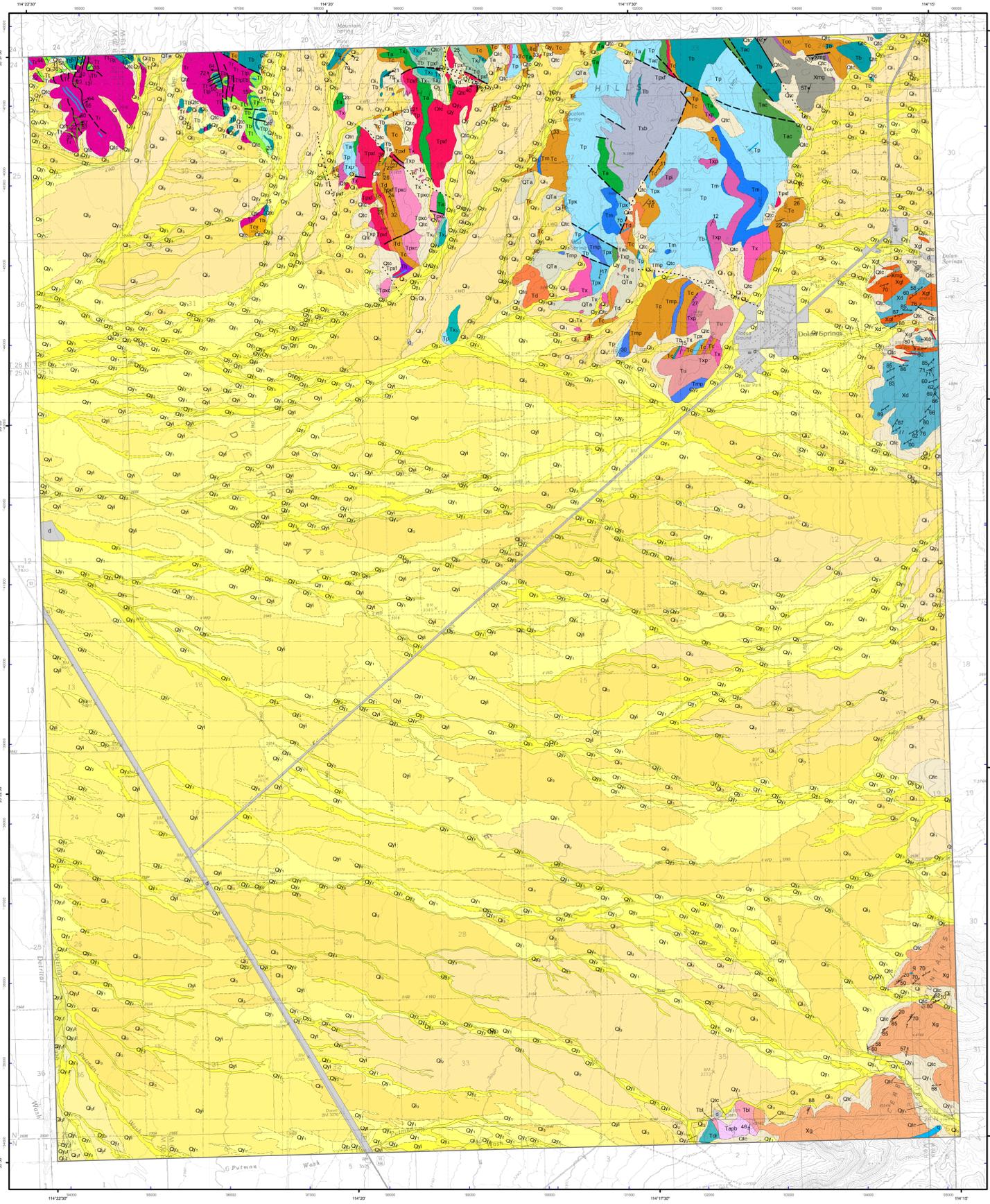
An apparently younger bimodal sequence of basalt-rhyolite occurs in the northwest corner of the  
map area that is separated from the other volcanics by a poorly understood west-dipping normal  
fault. The bimodal sequence is tilted into a series of north-striking, extensional synforms and  
antiforms.

## Stratigraphic Correlation Diagram



## Map Unit Descriptions

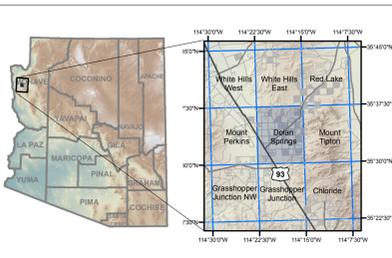
- Piedmont Deposits**
  - Qy** Modern channel alluvium - Unconsolidated, very poorly sorted sand, pebbles, cobbles and locally boulders in larger channels of piedmont washes. Channels may exhibit bar and swale microtopography with bars composed of coarser sediments. Cyclic deposits are typically unvegetated and exhibit no soil development although small shrubs and grasses can be found on slightly elevated bars and trees line some channel reaches.
  - Qy1** Latest Holocene alluvium - Poorly sorted, unconsolidated sand, pebbles, cobbles, silt and boulders located in small channels, on floodplains and on low-lying terraces. Vegetation cover and density is relatively large. Alluvial surfaces commonly exhibit bar and swale microtopography, but may be quite planar where deposits are fine. Soil development is generally absent on Qy1 surfaces. Qy1 surfaces are typically composed of the fluvial systems, and are susceptible to inundation during moderate to extreme flow conditions when flow exceeds channel capacity.
  - Qy2** Late Holocene alluvium - Poorly sorted, unconsolidated sand, pebbles, cobbles, silt and boulders located in small channels, on floodplains and on low-lying terraces. Vegetation cover and density is relatively large. Alluvial surfaces commonly exhibit bar and swale microtopography, but may be quite planar where deposits are fine. Soil development is generally absent on Qy2 surfaces. Qy2 surfaces are typically composed of the fluvial systems, and are susceptible to inundation during moderate to extreme flow conditions when flow exceeds channel capacity.
  - Qy3** Early to late Holocene alluvium - Poorly to very poorly sorted, weakly consolidated sand, silt, pebble, cobble and boulder deposits associated with low to intermediate terraces, abandoned drainages, and inactive areas. Fine and coarse sand and silt topography is common where deposits are gravelly. Soil development is weak, with minimal soil structure and visible carbonate accumulation, but gravelly surfaces appear darker brown than younger Qy2 surfaces due to moderate rock varnish accumulation.
  - Qy4** Holocene alluvium, undivided - Holocene alluvium, undivided
  - Qy5** Holocene to latest Pleistocene alluvium - Sand, silt, and clay with some pebble and cobble concentrations in low-relief alluvial fan / sheetflood areas. Rock varnish on surface gravel is weak to moderate. Surfaces are typically rounded, with some reddish brown, and soil development is weak to moderate, with some clay accumulation and stage II calcic horizon development. Washes typically are incised ~1 m below Qy1 surfaces.
  - Qy6** Late Pleistocene alluvium - Very poorly sorted, weakly consolidated sand, pebbles, cobbles, silt and some small boulders associated with inactive alluvial fans and higher terraces. Surfaces are typically dissected, with local topographic relief varying from about 1 to 10 m. Original depositional topography typically is modified by erosion. Surfaces are typically rounded, with some reddish brown, and soil development is weak to moderate, with some clay accumulation and stage II calcic horizon development. Washes typically are incised ~1 m below Qy1 surfaces.
  - Qy7** Middle to late Pleistocene alluvium - Very poorly sorted, weakly to moderately consolidated sand, pebbles, cobbles, silt, and clay deposits associated with dissected alluvial fans. Surfaces are drained by extensive, incised tributary drainage networks. Surfaces are typically dissected, with local topographic relief varying from about 1 to 10 m. Original depositional topography typically is modified by erosion. Surfaces are typically rounded, with some reddish brown, and soil development is weak to moderate, with some clay accumulation and stage II calcic horizon development. Washes typically are incised ~1 m below Qy1 surfaces.
  - Qy8** Early to middle Pleistocene alluvium - Very poorly sorted, weakly to moderately consolidated sand, pebbles, cobbles and boulders associated with deeply dissected alluvial fans. Locally surfaces are rounded, but more commonly they are light gray due to common calcic fragments. Soil development is variable, but weakly developed petrocalcic horizons (stage IV) are common. Surfaces are moderately to deeply dissected, with local topographic relief varying from about 2 to 10 m. Original depositional topography has been moderately modified by erosion, and surfaces are rounded and seldom planar. Surfaces are higher and are much more eroded than adjacent Qy1 surfaces.
  - Qy9** Middle to late Pleistocene alluvial deposits, undivided - Middle to late Pleistocene alluvial deposits, undivided
  - Qy10** Pliocene to early Pleistocene alluvium - Very poorly sorted, moderately consolidated cobbles, pebbles, boulders and sand associated with alluvial fan deposits underlying the highest alluvial ridges in the valley. Ridges typically are 5 to 10 m above modern active washes. Soil development is moderate and variable, and dominated by calcic carbonate accumulation. Surfaces are light to light color because they are covered with debris chummed up from indurated petrocalcic horizons and unvarnished to lightly varnished gravel.
- Axial Stream Deposits**
  - Qyr** Modern channel deposits - Unconsolidated, very poorly sorted sandy to cobbly beds exhibiting bar and channel microtopography. Clasts are typically subangular to subrounded, with mixed lithologies including fine-grained volcanics, granite and metamorphic rocks. Deposits are typically unvegetated to lightly vegetated and exhibit no soil development. These deposits are the first to become submerged during moderate to extreme flow events and can be subject to deep, high velocity flow and lateral bank erosion.
  - Qyr1** Flood channel and low terrace deposits - Sand, pebble, cobble and silt deposits in recently active flood channels and lightly vegetated in-channel bars and small planar flood terraces less than 1 m above active channel. These deposits do not exhibit soil development but commonly have a light vegetation cover of small trees and shrubs and grasses due to their relatively frequent inundation. These surfaces are commonly inundated under moderate to extreme flow events and can be subject to deep, high velocity flow and lateral bank erosion.
  - Qyr2** Latest Holocene to historical river deposits - Sand, silt, clay, pebbles and cobbles associated with low river terraces. In roughly the southern 1/2 of the quadrangle these terraces were probably the floodplain prior to historical channel entrenchment. Qy2 deposits are associated with broadly planar surfaces that locally retain the shape of the historical valley bottom. Qy2 surfaces are up to 2 m above modern Qy1 deposits and are the most extensive river terraces in the valley. Qy2 sediments were deposited when the Detrital Wash was widespread, stratiform-flowing river system and are dominated by fine grained floodplain deposits. These surfaces appear predominantly fine grained at the surface due in part to shallow flood inundation, but are composed of interfingering coarsely to pebbly braided channel and fine sand to silt floodplains. Where Qy2 deposits are moderately to deeply incised they are not subject to inundation by river floods, but they may be flood-prone in areas with less channel incision. Qy2 deposits are subject to catastrophic bank failure due to undercutting and lateral erosion during flow events.
  - Qyr3** Late to early Holocene terrace deposits - Sand, silt, pebble, clay and cobble deposits associated with low river terraces. In roughly the southern 1/2 of the quadrangle these terraces were probably the floodplain prior to historical channel entrenchment. Qy3 deposits are associated with broadly planar surfaces that locally retain the shape of the historical valley bottom. Qy3 surfaces are up to 2 m above modern Qy1 deposits and are the most extensive river terraces in the valley. Qy3 sediments were deposited when the Detrital Wash was widespread, stratiform-flowing river system and are dominated by fine grained floodplain deposits. These surfaces appear predominantly fine grained at the surface due in part to shallow flood inundation, but are composed of interfingering coarsely to pebbly braided channel and fine sand to silt floodplains. Where Qy3 deposits are moderately to deeply incised they are not subject to inundation by river floods, but they may be flood-prone in areas with less channel incision. Qy3 deposits are subject to catastrophic bank failure due to undercutting and lateral erosion during flow events.
  - Qyr4** Middle to late Pleistocene river terrace deposits - Sand, silt, clay, pebbles and cobbles associated with an older level basin floor. Deposits are compositionally similar to Qy3 deposits, but reflect alluvial surface occupy higher positions in the landscape. Terrace surfaces are slightly to moderately rounded. Vegetation is sparse, consisting of small shrubs and grasses. Soil development is moderate, with weak soil redening and stage II calcic horizon development. Qy4 surfaces are typically found as high-standing surfaces abutted by distal fan alluvium or as small remnants surrounded by younger alluvial valley deposits.
- Other Surficial Units**
  - d** Disturbed areas - Areas primarily disturbed by human activity. Road embankments, stockpiles, and mining disturbance.
  - Qc** Talus and colluvium - Angular to subangular, locally derived, weakly bedded to unbedded, moderately to steeply dipping slope deposits associated with bedrock hills and mountains. Most deposits are colluvium consisting of very poorly sorted cobbles, boulders, sand and clay. Local coarse framework deposits (talus) have little or no material at the surface.
- Tertiary rock units of the southern White Hills**
  - Yc** Younger conglomerate - A single exposure of mixed andesitic and felsic volcanic and Proterozoic gneissic clast conglomerate in the northwest part of the map area. The conglomerate is dominated by boulder-sized subangular to sub-rounded Proterozoic clasts and appears to overlie a sequence of east-titled felsic tuff and basaltic lava with angular unconformity.
  - Yr** Rhyolite lava - Rhyolitic lavas containing 3-15% 2-mm feldspar (typically diopside) and 1-3% 0.2mm biotite and hornblende. In the field database, some of the rhyolite flows are classified as phenocryst-rich (Yr1) if they contain over 12-15% phenocryst and some 1% difficult to evaluate. The content of the vitric portion of many flows, this classification was not used to define map units.
  - Yt** Felsic tuff - Nonwelded thin lapilli-rich, and locally pumice lapilli-rich ash-flow tuff and ash-fall tuff, typically thin to thick-bedded in tabular sets. The tuffs form fairly thick sequences interbedded with rhyolite lava flows that locally contain minor amounts of pumiceous sandstone and felsic volcanic dominated conglomerate.
  - Yb** Ash-flow tuff - Massive, poorly welded peach-colored felsic ash-flow tuff containing up to 40% phenocryst-poor to aphyric rhyolitic pumice lapilli and up to 7% 1-4mm plagioclase and sparse biotite phenocrysts. The ash-flow tuff forms an important marker unit in the northwest on an ridge where it is interbedded with nonwelded tuff (Yt) and basaltic lava (Yb).
- Tertiary rock units of the northwestern Cerbat Mountains**
  - Td** Phenocryst-rich dacitic and andesitic lava - Unit contains <2% 4mm biotite, 7-9% 2mm pyroxene, and <20% 4-8mm plagioclase. Pyroxene also occurs in sizes up to 6mm diameter. This unit, exposed near Larson Dam in the southeastern corner of the map area, is probably the same unit as mapped to the south in the Grasshopper Junction Quadrangle (map unit Td1 of Ferguson et al., 2009).
  - Tp** Pyroxene - biotite andesite - Medium to dark gray lava with ~1% 4mm biotite, ~2-4% 2mm pyroxene. Weathers reddish brown to black and is weakly vesicular to non-vesicular. This unit is exposed in the southeastern corner of the map area near Larson Dam.
  - Tb** Basalt of Larson Dam - Vesicular, black-weathering basalt near the southeastern corner of the map area. Unit contains 5-8% 2mm, black to greenish black pyroxene, 2-4% 2mm plagioclase, and 8-12% 1mm horn-blade spots inferred to be oxidized olivine. Red scoriaceous zones separate lava flows. Contact with granite base of basalt is covered over 2m interval but rocks on either side of the contact are not crushed or significantly fractured - the contact is inferred to be depositional.
- Proterozoic crystalline rocks**
  - Q** Vein quartz - Vein quartz over a large area in the southeastern corner of the map area consists of numerous, vertically anastomosing veins. Unit contains inclusions between veins of host granite that is foliated and appear to extend into 1.10-cm long iron-oxide rich mineral selvages within the vein quartz. Lack of deformation in vein quartz suggest that it is post-deformation and that it was emplaced along foliation surfaces, incorporating mineral selvages from host granite with veins. It is possible that vein quartz is older than foliation and that foliation within quartz is obscure and only visible in scoriaceous and selvages, but this alternative is less favored because the foliation in vein quartz was not observed.
  - Xf** Fine-grained biotite granite - Fine-grained biotite granite with highly variable mafic content (2-20%). This granite intrudes host coarse dioritic granitoid concordant to foliation and is a basal foliated. Weak tholitic layering is defined by quartz veins and vague compositional layering. The unit is recrystallized so that biotite is randomly oriented. Could be syngenetic with coarse dioritic granitoid as seems laterally associated and grades into fine-grained dioritic rock with 20% mafic minerals.
  - Xd** Coarse-grained dioritic granitoid - Dioritic granitoid consists of ~35% mafic minerals and ~65% feldspar and quartz, most or all of which is probably plagioclase. Mafic minerals consist of feldspar aggregates with highly undeformed minerals <1 mm diameter. Plagioclase is recrystallized and locally forms augen up to 2 cm long. This unit locally contains rounded, fine-grained biotite granitoid clasts that locally display this unit's assemblage as gneiss. Foliation in dioritic granitoid also affects pegmatites.
  - Xp** Foliated porphyritic granite - Recrystallized granite with spotted, relic pink K-feldspar up to 3 cm that are generally obliterated by strong recrystallization. Foliation is weak and inconsistent in orientation (away to swirl). High-temperature mylonitic foliation is apparent locally, but substantial mylonitization obscures mylonitic fabric. It is possible that mylonitization occurred at such high temperature that recrystallization almost kept pace with grain-size reduction during shearing.
  - Xm** Garnet gneiss - Quartz-biotite-garnet gneiss - Banded to laminated gneiss, with <1mm quartz and biotite, and 0-10% garnet up to 8 mm diameter. Unit is slightly porphyritic. Lithologic layering varies from sorted to regular to planar.



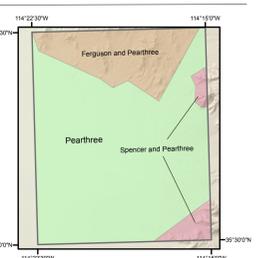
Topographic base from USGS 1:24,000 Quadrangle Series.  
Seamless basemap generated using iCage All Topo Pro software.  
Hillshade generated from NED 10m Digital Elevation Model.

Projection Information:  
North American Datum of 1983  
1000-meter Universal Transverse Mercator grid tics, zone 11N, shown in blue.

## Location Map



## Mapping Responsibility



## Map Symbol Descriptions

- Tectonic foliation defined by preferred mineral orientation with or without compositional layering. This fabric indicates strongly annealed mylonitic foliation.**
- Mylonitic foliation, slightly to moderately annealed.**
- Bedding, inclined**
- Generic foliation defined by preferred mineral orientation with or without compositional layering.**
- Generic lineation**
- Slickenside striae lineation**
- Dip of fault, dike, or contact**
- Cumulate foliation**
- Contact, accurate**
- Contact, approximate**
- Fault, approximate**
- Fault, accurate**
- Intrufonnational Contact**
- Generic foliation, inclined**
- Generic lineation**
- Slickenside striae lineation**
- Dip of fault, dike, or contact**
- Cumulate foliation**
- Fault, concealed**
- Quart vein**
- Mafic dike**
- Debris-flow levee**

## References

Ferguson, C.A., Johnson, B.J., Pearthree, P.A., and Spencer, J.E., 2009. Geologic map of the Grasshopper Junction 7 1/2 Quadrangle, Mohave County, Arizona, v. 1: Arizona Geological Survey Digital Geologic Map DGM-70, 2 sheets, scale 1:24,000.

Mason, D.A., Ivanic, P.A., Conway, B.D., Kurtz, J.A., and Winn, M.T., 2007. Preliminary estimate of ground water in storage for the Detrital Valley ground-water basin, Mohave County, Arizona: Arizona Dept. of Water Resources Open-File Report Number 9, 31 p., 4 Appendices.

Richard, S.M., Shipman, T.C., Greene, L.C., and Harris, R.C., 2007. Estimated depth to bedrock in Arizona v. 1.0: Arizona Geological Survey Digital Geologic Map DGM 52, 9 p., text, scale 1:1,000,000.